



Simulations of a Convective Boundary Layer with a Dynamic Smagorinsky Scheme

Modelling the Greyzone Boundary Layers (GreyBLs)

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SIR Modelling Turbulence with grid spacing (Δ)

BL resolved

•∆ of O~50m ·Used for research purposes ·Assumption: $\Delta < < \ell$ Resolves large eddies ·Small eddies parametrized

Greyzones

 $\cdot\Delta$ of O~1km Starting to be used in **NWP** ·Situation: ∆≈ℓ ·LEM and low resolution parametrizations not suitable

BL not resolved

 $\cdot \Delta > 1 \text{km}$ ·Used over many years for weather forecasting and climate modelling ·Assumption: Δ>>ℓ ·BL- parametrized fully

Λ

- Convective Boundary Layer in grey zones because large eddies scale with boundary layer height
- Dynamic model considered as one possible model suitable for greyzones.
- Stretch grid length towards coarse resolution to determine point where the subgrid model choice makes a difference.

Smagorinsky-Lilly in LEM

$$\frac{Du_i}{Dt} = -\frac{\partial}{\partial x_i} \left(\frac{p'}{\rho_s}\right) + \delta_{i3}B' + \frac{1}{\rho_s}\frac{\partial \tau_{ij}}{\partial x} - 2\epsilon_{ijk}\Omega_j u_k$$

 $\tau_{ij} = \rho_s \lambda^2 f_m \left(R i_p \right) S S_{ij}$

$$S_{ij} = \frac{\partial u_i}{\partial x_j} + \frac{\partial u_j}{\partial x_i}$$

$$S = \left(\frac{1}{2} \sum_{i,j=1,3} S_{ij}^2\right)^{1/2}$$
$$\frac{1}{\lambda^2} = \frac{1}{\lambda_0^2} + \frac{1}{[k(z+z_0)]^2}$$



 $\lambda_0 = c_s \Delta$





Dynamic Model

- Number of studies determined c_s
 - Flow dependent and suggested values include 0.1, 0.2 and 0.23.
 - Germano 1991 introduced a method that allows c_s to be determined from the flow $G_{\alpha\Delta}^*(...)$

$$\tau_{ij} = -2c_{s,\Delta}^2 \Delta^2 |\tilde{S}| \tilde{S}_{ij}$$

$$T_{ij} = -2c_{s,\alpha\Delta}^2 \left(\alpha\Delta\right)^2 |\overline{S}| \overline{S_{ij}}$$

$$L_{ij} = \overline{\tilde{u}_i \tilde{u}_j} - \overline{u}_i \overline{u}_j$$

$$L_{ij} - \left(\overline{T_{ij}} - \overline{\tau_{ij}}\right) = e = L_{ij} - c_{s,\Delta}^2 M_{ij}$$

$$M_{ij} = -2\Delta^2 \left(\overline{|\tilde{S}|\tilde{S}_{ij}} - \alpha^2 \beta \overline{|\overline{S}|\overline{S}_{ij}} \right), \beta = c_{s,\alpha\Delta}^2 / c_{s,\Delta}^2$$









- Germano (1991) plane averaged
 - Suitable for horizontally homogeneous flows
- Meneveau et al (1996) Lagrangian averages
 - Suitable for inhomogeneous flows and complex geometries
- Bou-Zeid et al. (2005) Lagrangian averaged scale variant
 - Uses a second test scale to determine $\beta = C_{s,4\Delta}^2 / C_{s,2\Delta}^2$
 - Proposed as a procedure that could be suitable for the grey zone



Simulations



- Met Office Large Eddy Model
- Convection atmosphere
 - Constant sensible heat flux : 241Wm⁻²
 - Constant temp of 300K up to 1km, and a sharp jump of 8K is imposed over a depth of 100m near the top of the BL.
 - 1K amplitude perturbations, 4 hour simulations
 - Weak geostrophic winds $(U_g, V_g) = (1,0)$ m/s
- Domain size 9600x9600x2000m

Δx	Δz=0.4*Δx	∧ ₀ =0.23*∆x	Grid points	
25	10	5.75	384x384x200	Only Smag
50	20	11.5	192x192x100	
100	40	23.	96x96x50	
200	80	18.4	48x48x25	
400	160	36.8	24x24x13	





- Plane-averaged scale invariant (PASI)
- Lagrangianaveraged scaleinvariant (LASI)
- Lagrangianaveraged scaledependent (LASD)





CS probability distribution





- 200 m resolution
- As resolution is decreased more clipping occurs



Resolved Potential temperature flux





- Decrease with height to a minimum
- Negative region = entrainment zone
- Minimum lower height with low resolution
- CG data is more converged below zi



Structures at different heights





- Thermals rise
- Join those in adjacent regions to form larger structures.
- Theta' gets smaller due to mixing
- Closer to BL height – negative theta' associated with positive w'



Temperature flux Quadrant analysis





- Disentangle the temperature flux.
- e.g.Sullivan et al., 1998; Coceal et al 2007, Park and Baik 2014
- Theta'>0,W'>0 : Q1
- Theta'<0,W'>0 : Q2
- Theta'<0,W'<0 : Q3
- Theta'>0,W'<0 : Q4
- Number of events and contribution of each quadrant to the total flux.



25m resolution Quadrant analyses





- Thermals rise mix with environment-get colder
- Some join Q3 closer to the surface, most become Q2
- Q2 bigger contribution theta'w'
- More Q4 events close to inversion layer- entrainment – contribution is about ¹/₄.
- Q4 events mix with the environmental air- become cold - Q3

Grid size comparison for Smag





- NO evidence of large entrainment with low resolution
- Q2

 contribution
 much larger
 than the rest
 with
 increased
 grid length



Contribution at 100m resolution – no f(Ri)





- Similar at higher resolution.
- More similar with stability functions.



W' at Z=60m deltax=50





- Mean variables are almost the same.
- Structures look different
- LASI is more broken – too noisy
- Smag is too smooth



500

250

0└ 0.1

0.2

0.3

200m Quadrant 3 Contribution

0.4

Qaudrant analyses – 200 m no f(Ri)







 Lines more divergent than without stability functions.



Z=160m Smag deltax=200m





- Smag and PASI are smooth.
- Lagrangian
 models look
 slightly better



Z=1120 deltax=200m

4000





- Strong w' in all subgrid models.
- Lagrangian models make no improvement here



Contribution at 400m resolution





Run without stability functions crashes

Simulatio
 ns much
 more
 divergent.





Summary

- Stability functions make a large difference from 200m and higher.
- Subgrid model choice make a small difference from 200m grid spacing and higher.
- No subgrid choice is better, or worse in the Convective BL when considering domain means.
- Improvements were found using the dynamic model for the morning transition (in progress).





Thank you for your attention